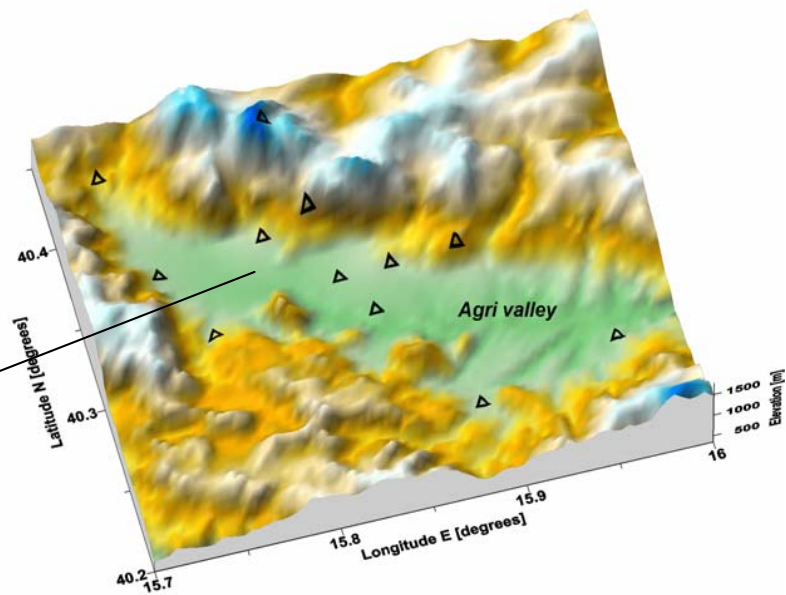
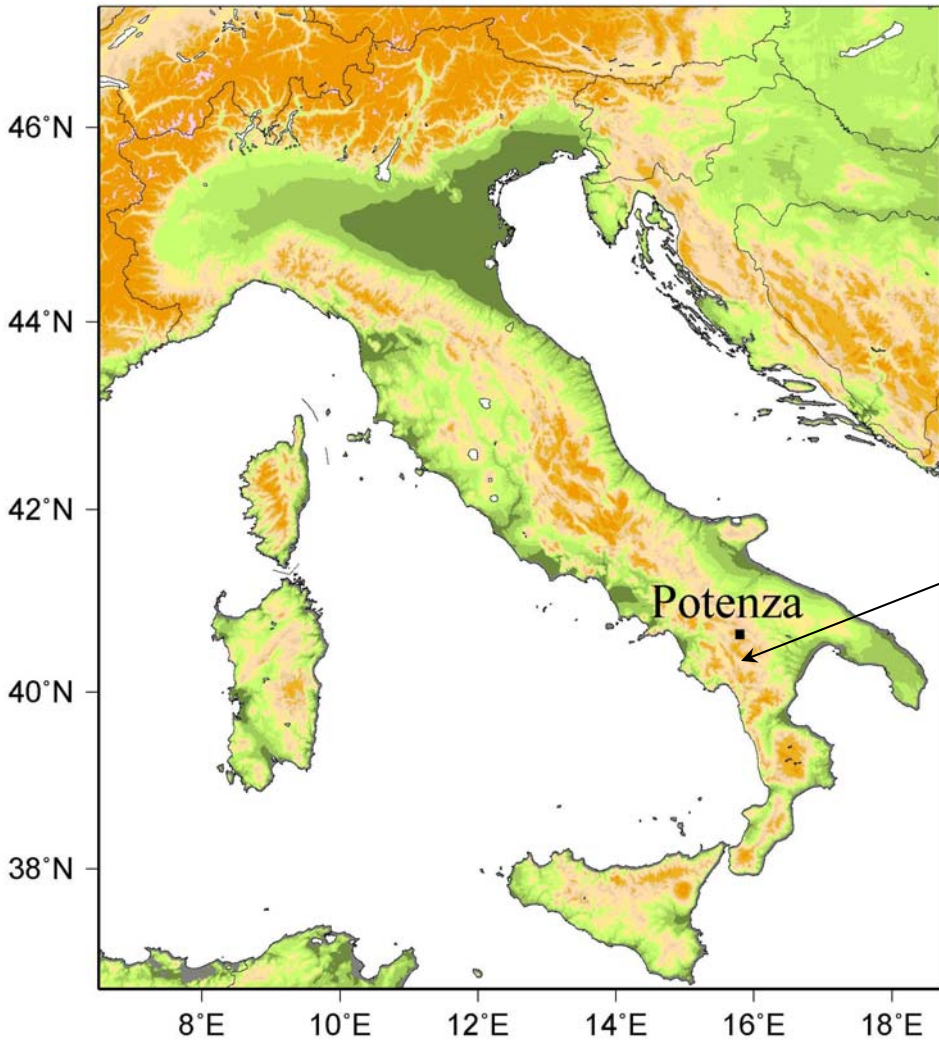


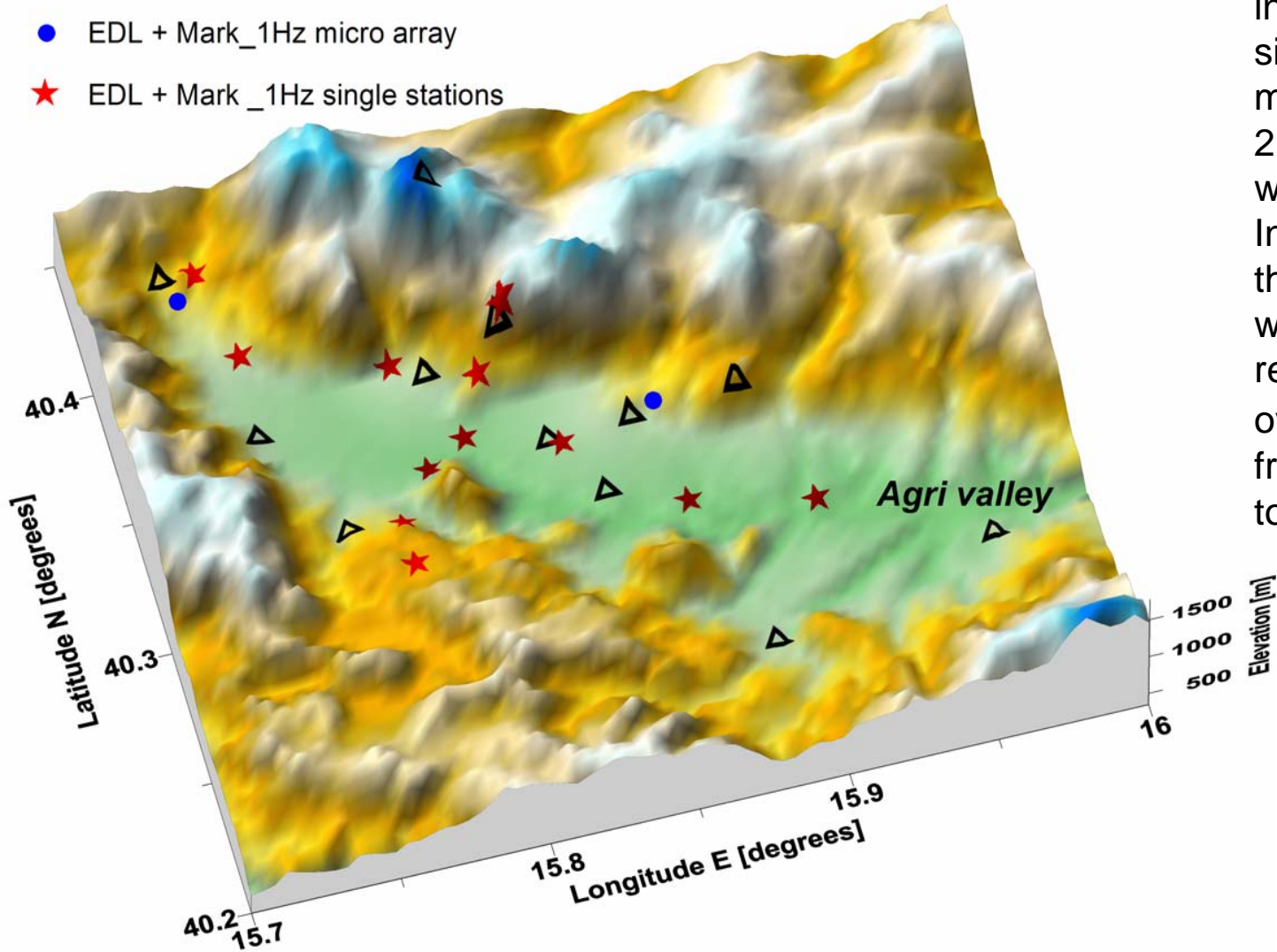
DATA



Agri Valley

DATA

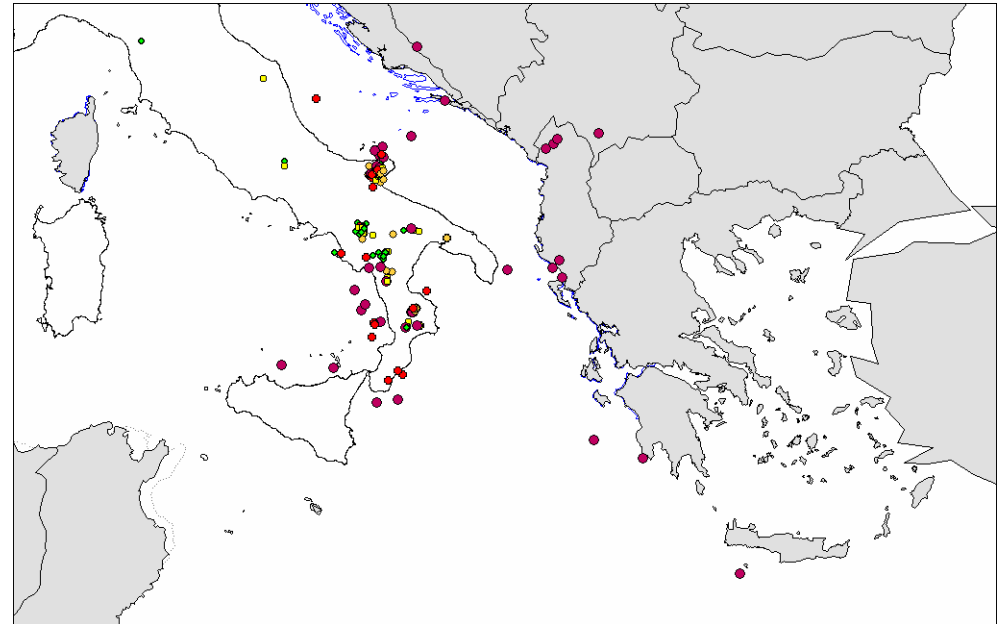
- △ EDL + Mark_1Hz array 1 month
- EDL + Mark_1Hz micro array
- ★ EDL + Mark_1Hz single stations



As preliminary investigation 12 single station measurements and 2 micro arrays were performed. In the mean time the 12 stations were installed to record earthquakes over ~ 1 month from 25/09/2006 to 25/10/2006

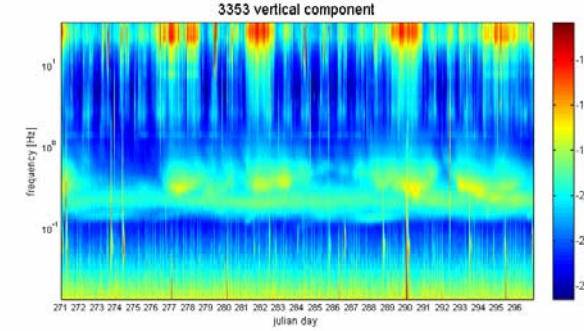
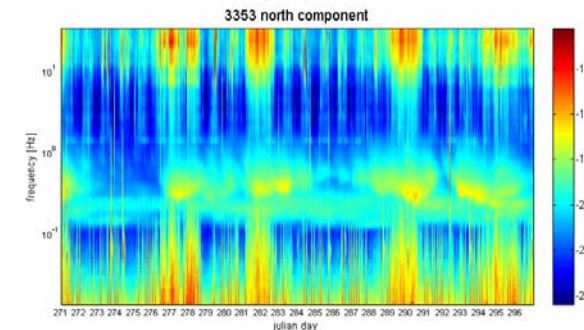
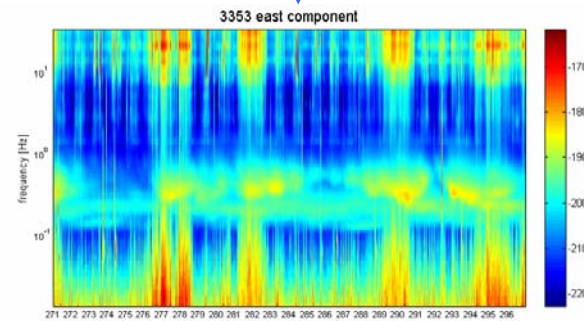
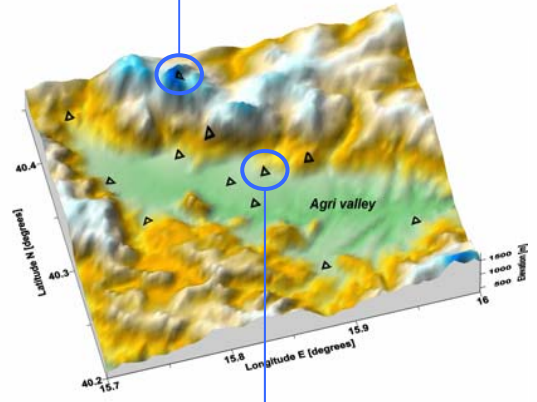
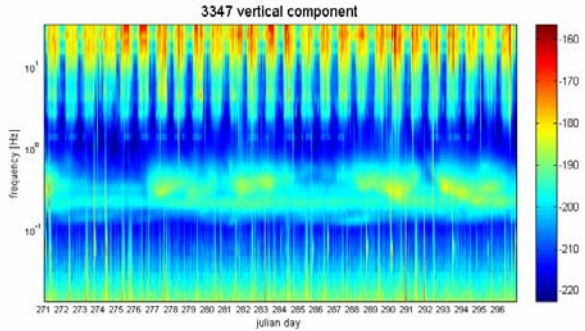
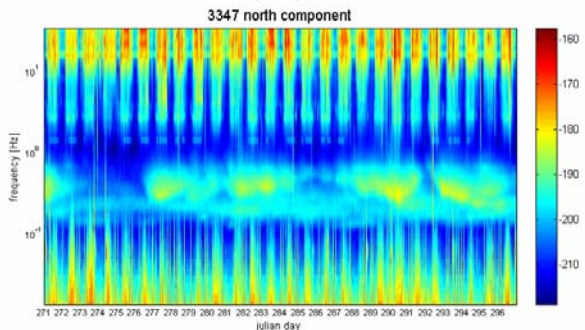
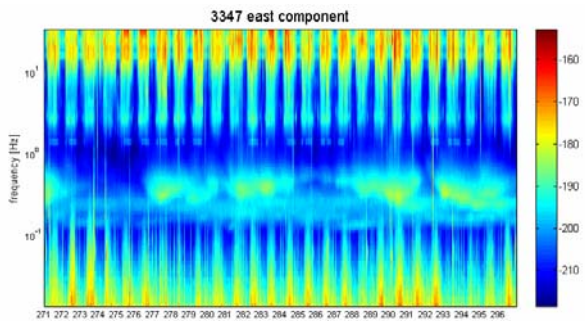
DATA

Month	Day	Hour	Epicentral area
Sep	25	9.36.00	Isole Tremiti
Sep	26	12.00.00	Costa calabra occidentale
Sep	26	2.24.00	Irpinia
Sep	30	14.24.00	Sud Grecia
Oct	4	12.00.00	Isole Tremiti
Oct	5	14.24.00	Appennino Lucano
Oct	8	19.12.00	Irpinia
Oct	14	19.12.00	Val d'Agri
Oct	14	12.00.00	Val d'Agri
Oct	14	2.24.00	Val d'Agri
Oct	14	14.24.00	Irpinia
Oct	15	4.48.00	Val d'Agri
Oct	15	19.12.00	Val d'Agri
Oct	16	16.48.00	Adriatico meridionale
Oct	16	9.36.00	La Sila
Oct	16	12.00.00	Val d'Agri
Oct	20	7.12.00	Cilento
Oct	22	12.00.00	Monte Alpi-Sirino
Oct	25	2.24.00	Irpinia
Oct	25	16.48.00	Irpinia



~ 100 events recorded in ~ one month.
 The 21 events listed in the table have
 a signal to noise ratio greater than 3

DATA



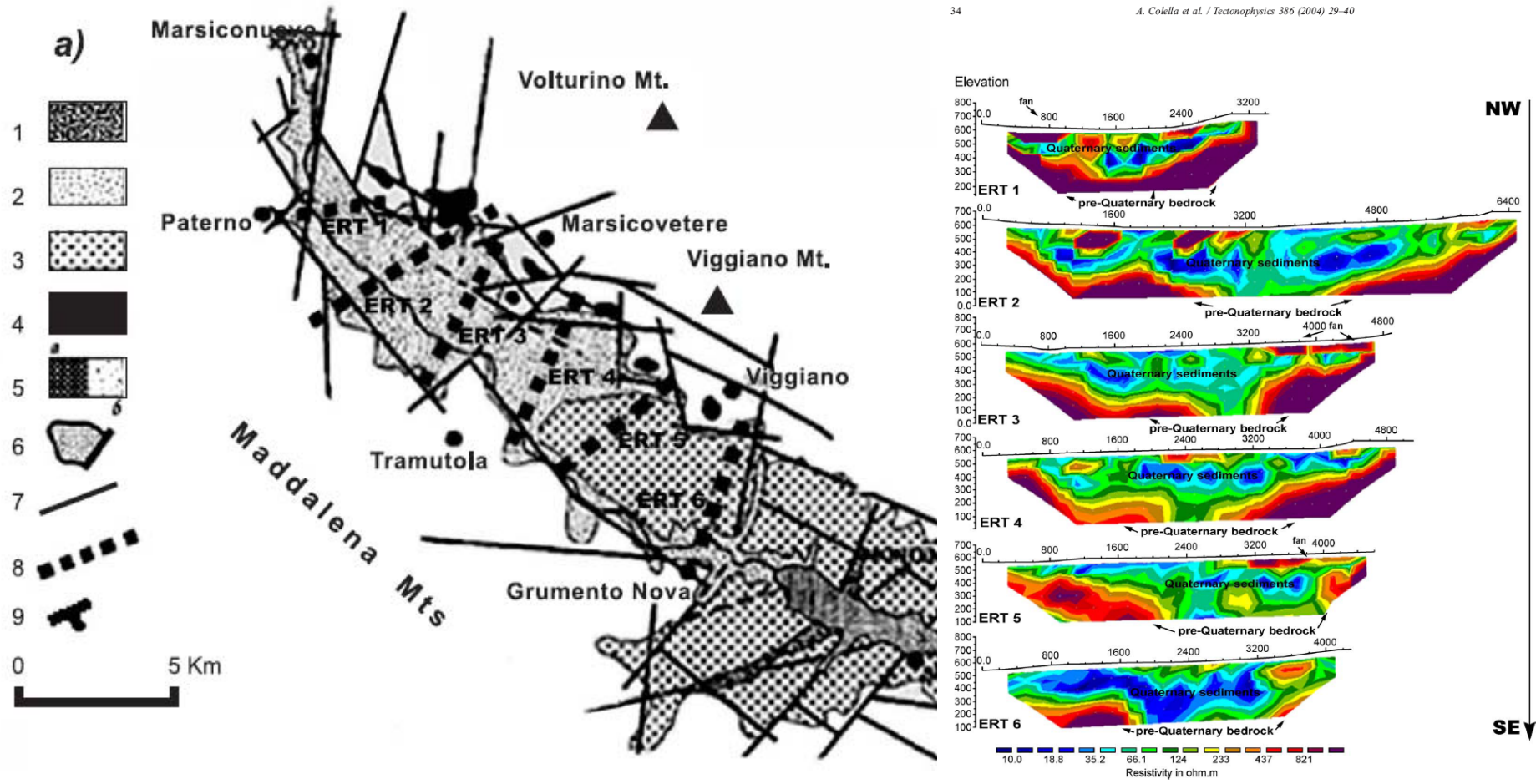


Fig. 1. (a, b) Simplified geological map of the Quaternary deposits of the High Agri Valley Basin with location of major faults and ERT profiles (dotted lines). Legend: (1) recent alluvial deposits; (2) alluvial deposits (Holocene–upper Pleistocene); (3) alluvial deposits (“Compleso Val d’Agri”, upper–middle Pleistocene); (4) subaerial slope deposits (middle–lower Pleistocene); (5) conglomerates and sands of the Sant’Arcangelo basin (Serra Corneta Fm [a] and Castronuovo Fm [b], lower Pleistocene); (6) Pietra del Pertusillo dam; (7) faults; (8) ERT surveys; (9) master fault located in c. Pre-Quaternary units are not shown (from Giano et al., 2000b). (c) Localization of master fault associated with the 1857 event (from Benedetti et al., 1998).

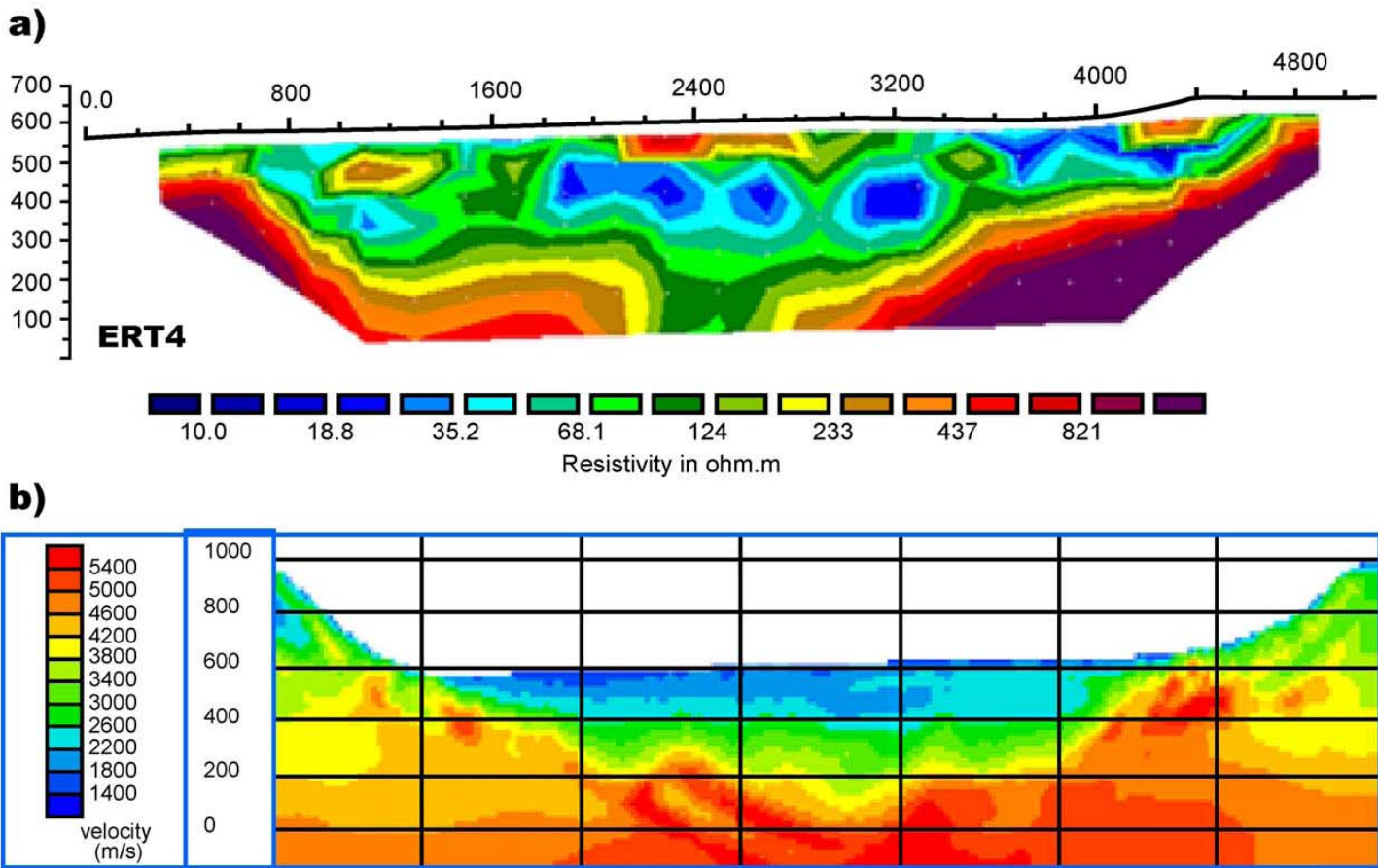


Fig. 4. Comparison between electrical resistivity tomography and seismic tomography. In Fig. 5a, the ERT-4 has been displayed; in Fig. 5b, a seismic tomography carried out along the same profile by Enterprise Company (Dell’Aversana and Morandi, 2002) is reported. The increase of the seismic velocity and resistivity with the depth of the investigated layers is quite similar.

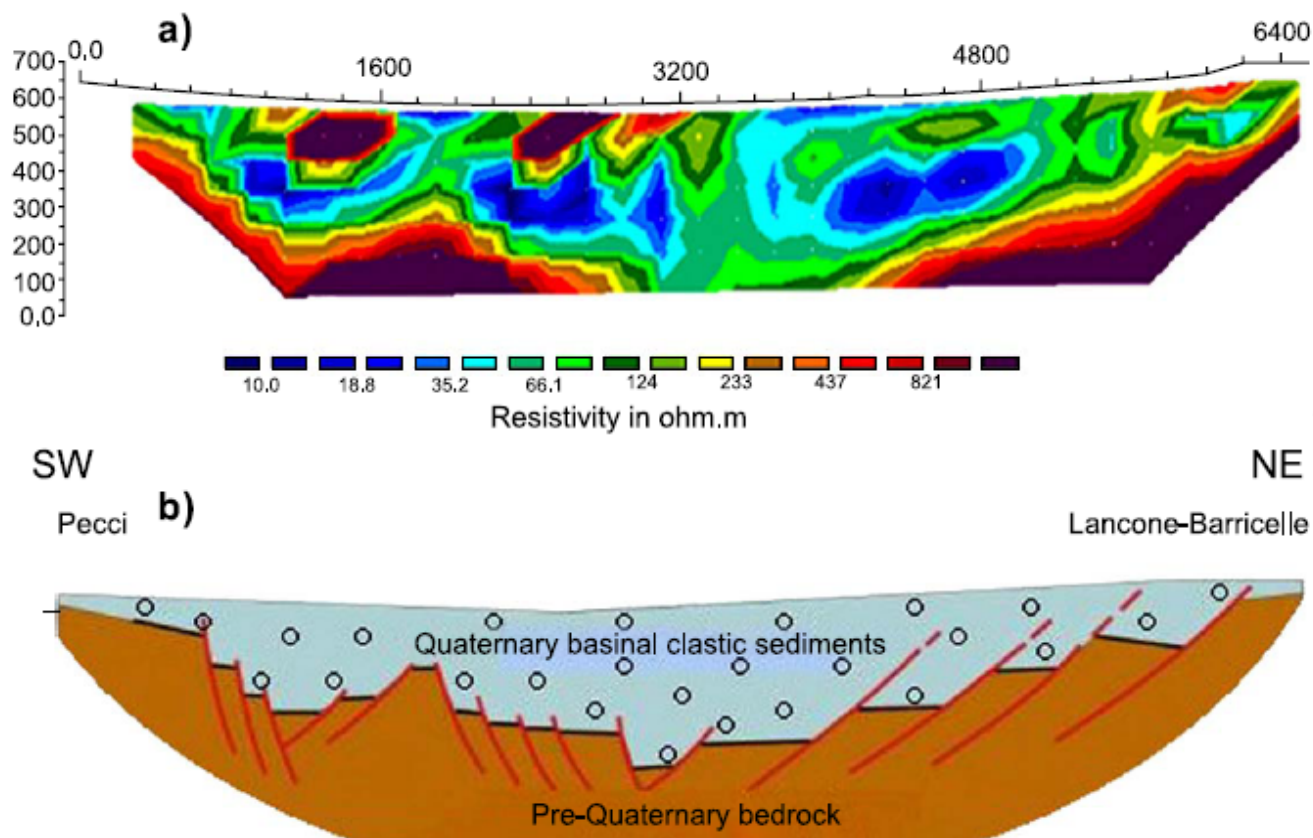
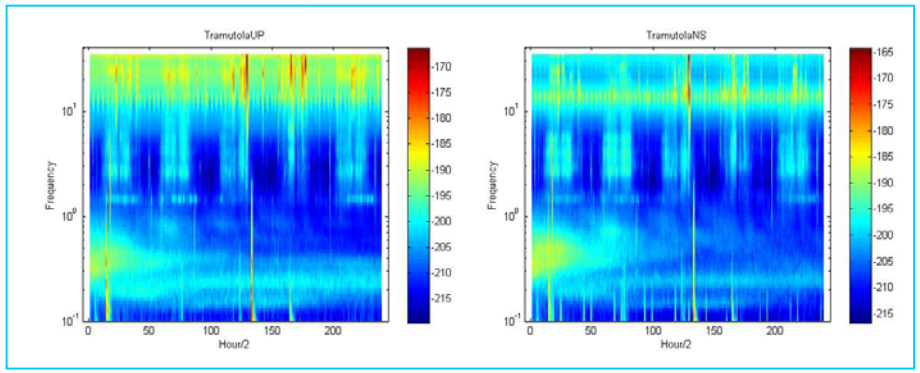
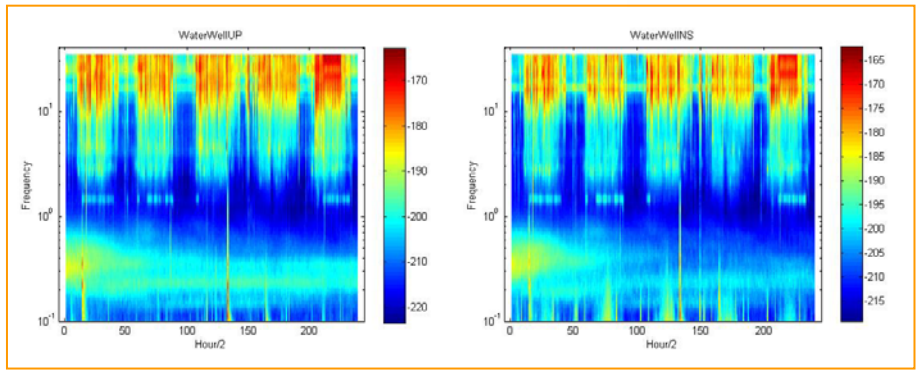
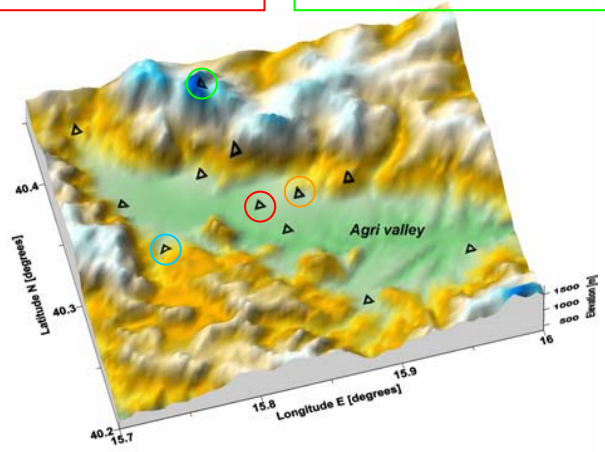
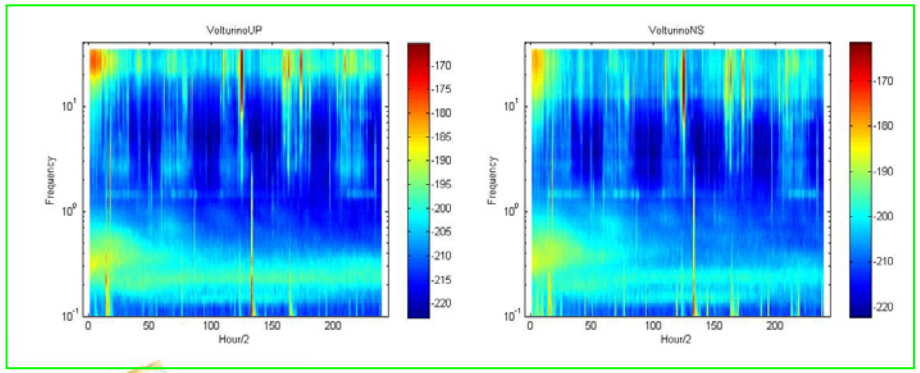
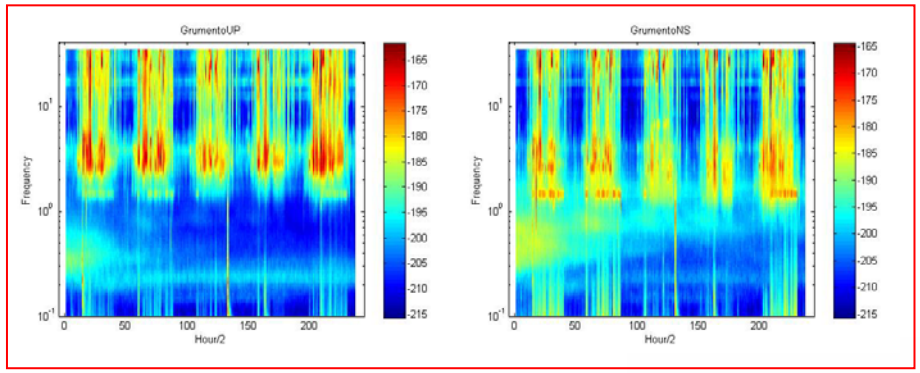
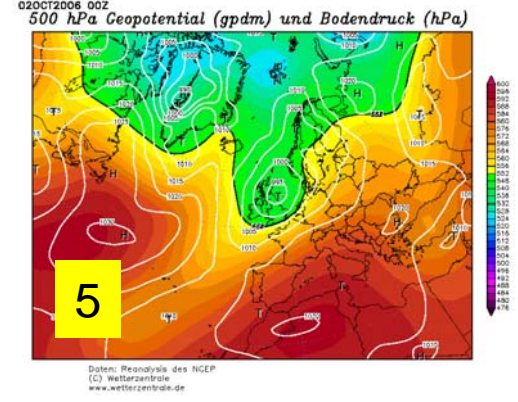
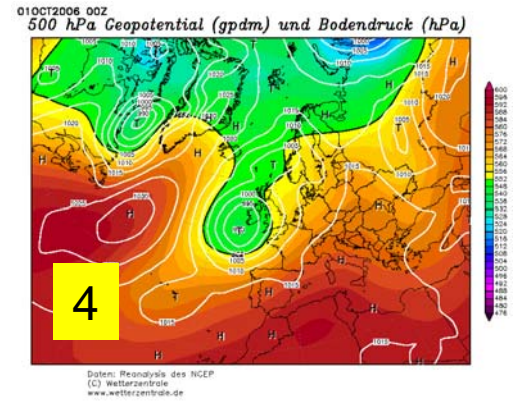
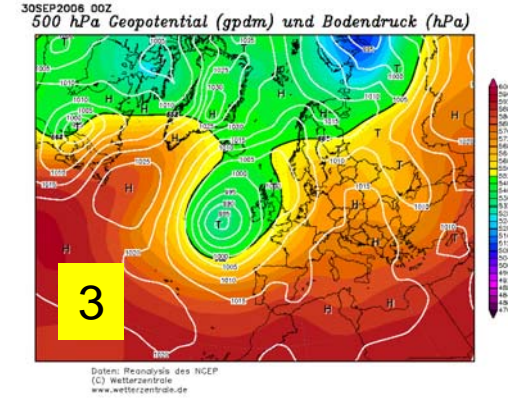
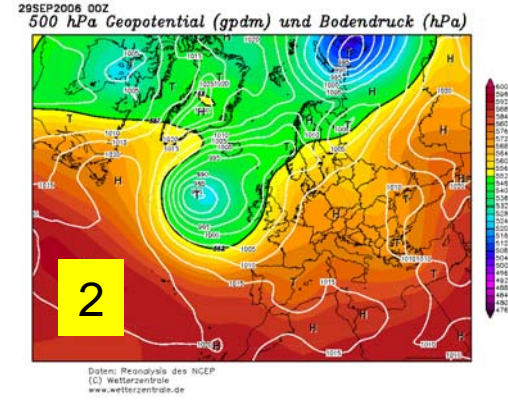
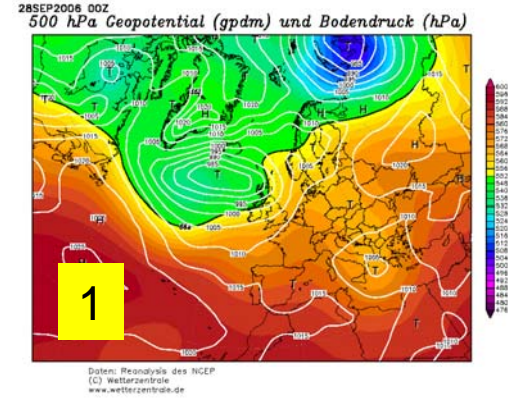
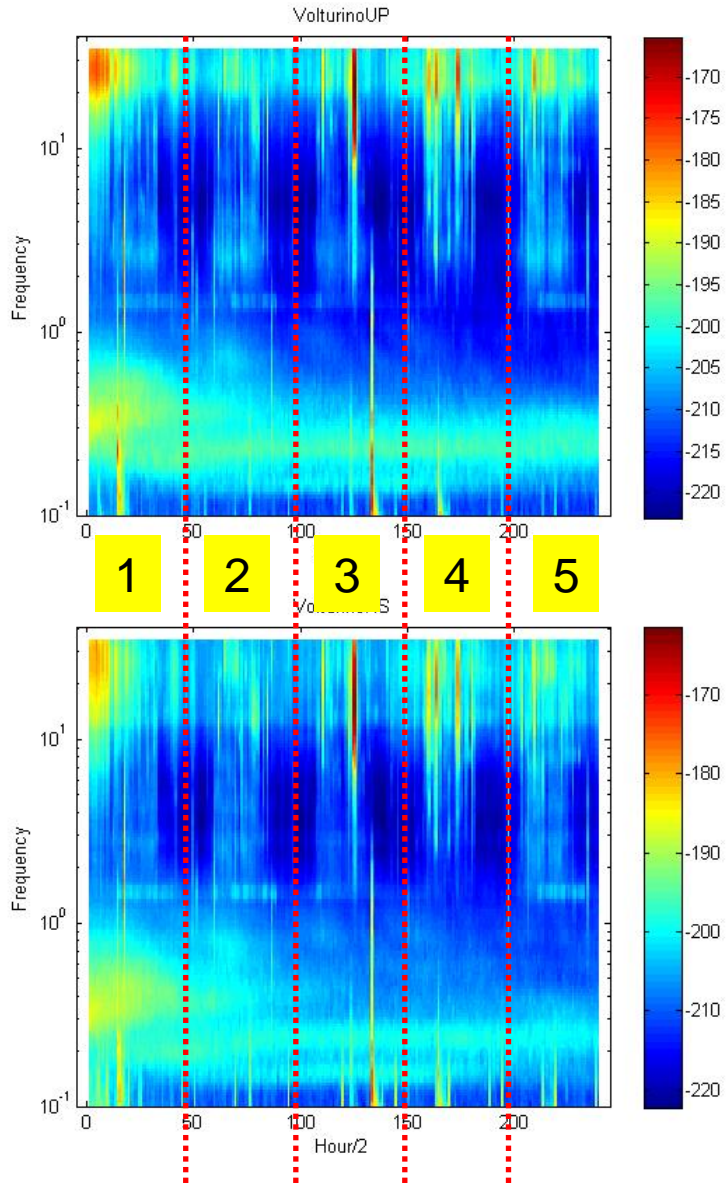


Fig. 5. The figure shows a simplified geological section transversely to the northern part of the High Agri Valley. It marks the contact between the upper clastic basin fill and the lower pre-Quaternary bedrock, and the synsedimentary faults could dissect it. The basin, up to 500 m deep, could appear as an irregular graben with a minor graben due to antithetic faults.

Seismic Noise



Seismic Noise



Seismic Noise

